

## Prediction of weathering development in metarhyolites of the Ilgın (Konya) area, SW Turkey

Erkan BOZKURTOĞLU\*, Şenel ÖZDAMAR, Hatice ÜNAL ERCAN

Department of Geological Engineering, Faculty of Mines, İstanbul Technical University, 34469 Maslak, İstanbul, Turkey

Received: 04.04.2012 • Accepted: 14.07.2012 • Published Online: 27.02.2013 • Printed: 27.03.2013

**Abstract:** Fresh to weathered metarhyolites crop out in the Ilgın (Konya) area of the Afyon-Bolkardağ Zone. Determination of the development of weathering was studied by physical (i.e. specific gravity, dry unit weight, saturated unit weight, porosity, void ratio, and degree of saturation by weight) and mechanical (i.e. point load) properties and the “point rock change value” (RCV<sub>p</sub>) and “point rock change ratio” (RCR<sub>p</sub>) values of the metarhyolite rock samples. The samples were classified in 3 groups (i.e. A, B, and C) representing degree of weathering from weathered to fresh rocks based on their RCV<sub>p</sub> and RCR<sub>p</sub> values. The K<sub>2</sub>O values are 7.09 wt.%, 8.62 wt.%, and 8.75 wt.% and the matrix ratios are 60%-70%, 50%-60%, and 20%-25% for groups A, B, and C, respectively. The RCV<sub>p</sub> and RCR<sub>p</sub> values of the studied samples range between 0.952 and 0.99 and 4.973% and 0.989%, respectively. Calculations show that metarhyolites will be completely changed by weathering at a 9.01% RCR<sub>p</sub> value according to metarhyolite alkali values varying in the 8.12%-9.40% range, with the average value being 8.89%. At the end of the rock change processes by weathering, the rocks remain chemically as metarhyolite, while their physico-mechanical properties and mineralogical compositions change to become soil. The average K-Ar ages vary between 60.4 ± 0.9 Ma and 64.1 ± 2.00 Ma. The whole-rock alteration can furthermore be predicted by the relationships between the RCR<sub>p</sub> and K-Ar ages of the 3 groups, which indicate that the rocks will be fully altered in the next 4.593 and 9.393 Ma. The whole-rock alteration will be completed for group A rocks in 4.6 Ma, for group B rocks in 7.2 Ma, and group C rocks in 9.4 Ma, provided that all the weathering agents take effect under the same conditions across the area.

**Key Words:** Ilgın, metarhyolite, physico-mechanical properties, weathering, point rock change value, point rock change ratio, K-Ar ages, whole-rock alteration

### 1. Introduction

The study area is located north of the town of Ilgın (Konya Province) in the Afyon-Bolkardağ Zone (ABZ) (Figure 1). The geology, petrography, geochemistry, and K-Ar ages of the metamorphic rocks of the Ilgın area in the ABZ were described in detail by Özdamar *et al.* (2012). These rocks have various degrees of weathering features.

Weathering is the breakdown of rocks and minerals at and below the earth's surface by physical and chemical processes. The reaction of various agents with rocks and weathering processes are shown by changes in the mineralogical, chemical, physical, and mechanical properties and grain size, or alteration in the weathered material compared to the fresh rock. The changes produced in the fresh rock by weathering can be ascribed to partial or complete decomposition of some minerals, the stability of other minerals, the oxidation of ferrous (Fe<sup>2+</sup>) to ferric (Fe<sup>3+</sup>) iron, and the partial or complete mobilisation of both major and minor chemical elements (Carroll 1970). Consequently, the weight changes in rock by weathering and alteration processes are reflected as changes in physical

and mechanical properties of rocks. These changes can be measured and the discrete effects of weathering or chemical alteration, or both, may be explained with conventional methods for engineering purposes (i.e. Bell 1994). Two of the useful tools in determining the final rock condition are “point rock change value” (RCV<sub>p</sub>) and “point rock change ratio” (RCR<sub>p</sub>), proposed by Bozkurtoğlu (2003) and Bozkurtoğlu *et al.* (2006). RCV<sub>p</sub> represents the final physical, chemical, mineralogical, petrological, and mechanical conditions of rocks. RCV<sub>p</sub> values range from 1 to 0, where the rock condition goes from fresh to fully altered. RCR<sub>p</sub> is the rock change ratio related to RCV<sub>p</sub> values.

This study focuses on the examination of physical, mechanical, chemical, and mineralogical properties of weathered rocks in the Ilgın (Konya) area. The results allow us to predict the full change in rocks by weathering in metarhyolites in the Ilgın area using RCV<sub>p</sub> and RCR<sub>p</sub> with K-Ar age values, and this is the first study comparing these values.

\* Correspondence: erkan@itu.edu.tr

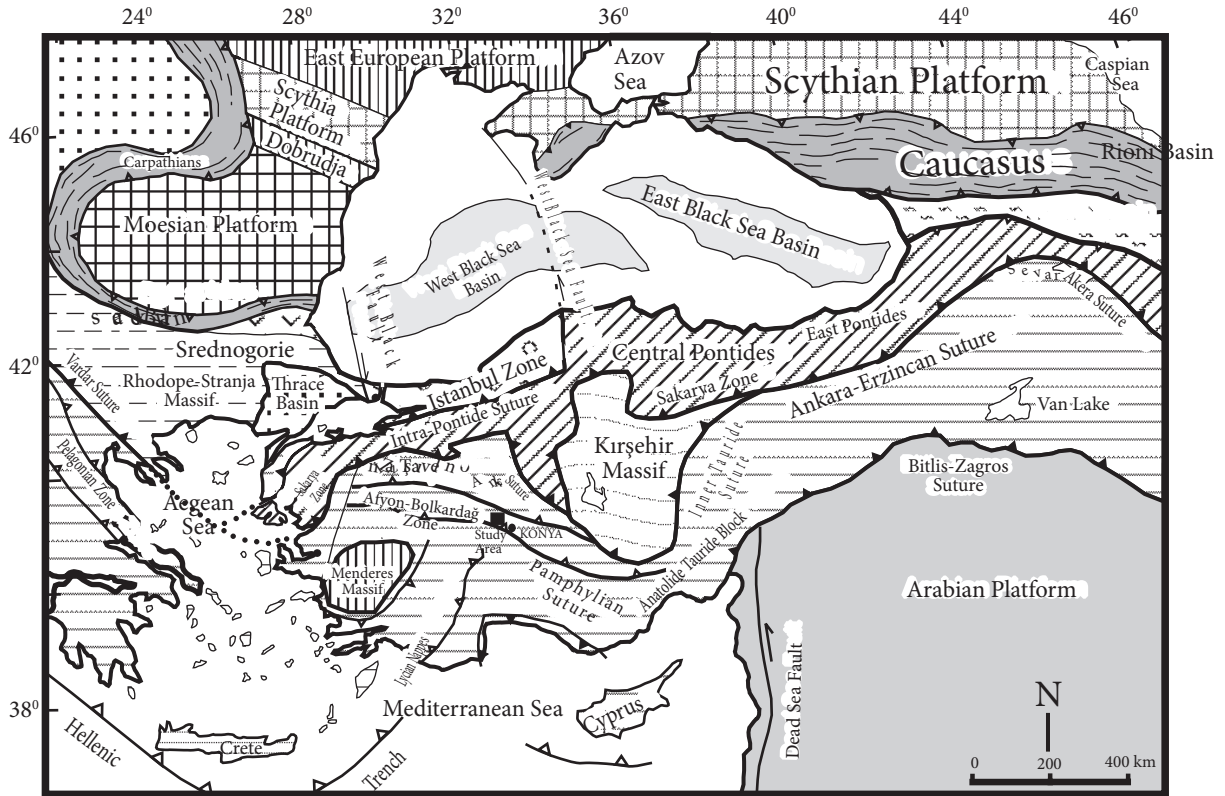


Figure 1. Tectonic units of the eastern Mediterranean Sea-Black Sea region (simplified after Okay & Tüysüz 1999).

## 2. Geology, petrography, geochemistry, and K-Ar, Ar-Ar, and U-Pb age dating

Two main metamorphic sequences, a Palaeozoic sequence and a Mesozoic sequence, which include metarhyolites, are unconformably overlain by Neogene cover with Quaternary alluvium in the Iğın (Konya) area, which is a province in the ABZ (Özdamar *et al.* 2012) (Figure 2).

The Palaeozoic metamorphic sequence contains metamorphosed conglomerate, sandstone, siltstone, claystone, limestone, and orthoquartzite. The Mesozoic metamorphic sequence consists of metaconglomerate at the base and fine-grained metasediments, metacarbonate, and intercalated metalavas and metatuffs at the top. Metarhyolites are mostly metalavas and unmapped metatuffs, which occur as thin beds within schists. Neogene sediments consist of yellowish and reddish conglomerate, sandstone, claystone, unconsolidated fragments, and locally carbonate-rich levels. The composition of metarhyolites is made up of 75%-80% groundmass and 20%-25% phenocrysts represented by quartz (Qtz), K-feldspar (Kfs), relict albite (Ab), and possibly sanidine (San). The matrix consists of fine-grained Qtz, Kfs, and Ab and newly formed extensive phengitic white mica. Accessory phases are zircon, rutile, epidote, and apatite (Özdamar *et al.* 2012). The chemical compositions of the metarhyolites are presented in Table 1.

The metarhyolites have 66%-77% SiO<sub>2</sub>, 12%-18% Al<sub>2</sub>O<sub>3</sub>, 5.8%-10.7% K<sub>2</sub>O, 0.07%-1.77% Na<sub>2</sub>O, 0.1%-1.1% MgO, and <1% CaO, and they plot in the rhyolite, comendite-pantellerite, or rhyodacite-dacite fields in the SiO<sub>2</sub> vs. Zr/TiO<sub>2</sub> diagram of Winchester & Floyd (1977) (Figure 3). Moreover, all samples except one are subalkaline in character (Figure 4) based on the classification of Irvine & Baragar (1971).

The K-Ar ages obtained from the whole-rock samples of metarhyolites are 60.4 ± 0.9 Ma, 62.6 ± 0.9 Ma, and 64.01 ± 2.0 Ma (Özdamar *et al.* 2012). The Ar-Ar phengite ages of the metarhyolites are 63.73 ± 0.06 Ma and 62.64 ± 0.12 Ma, and U-Pb zircon ages of the metarhyolites are 230 ± 2 Ma and 229 ± 2 Ma (Özdamar 2011).

## 3. Method for generating RCV<sub>p</sub> and RCR<sub>p</sub>

Specific gravity ( $\gamma_s$ ) is a critical measure of rock weathering and alteration (Browne 1998). This is measured by using a pycnometer and can be calculated from the phase diagrams of soil and rock. The ratio of measured values versus values calculated by specific gravity for each sample is the RCV<sub>p</sub> (Bozkurtoğlu 2003; Bozkurtoğlu *et al.* 2006). In phase diagrams of soil and rock, the relationship between dry unit weight ( $\gamma_d$ ) and specific gravity ( $\gamma_s$ ) is given by:

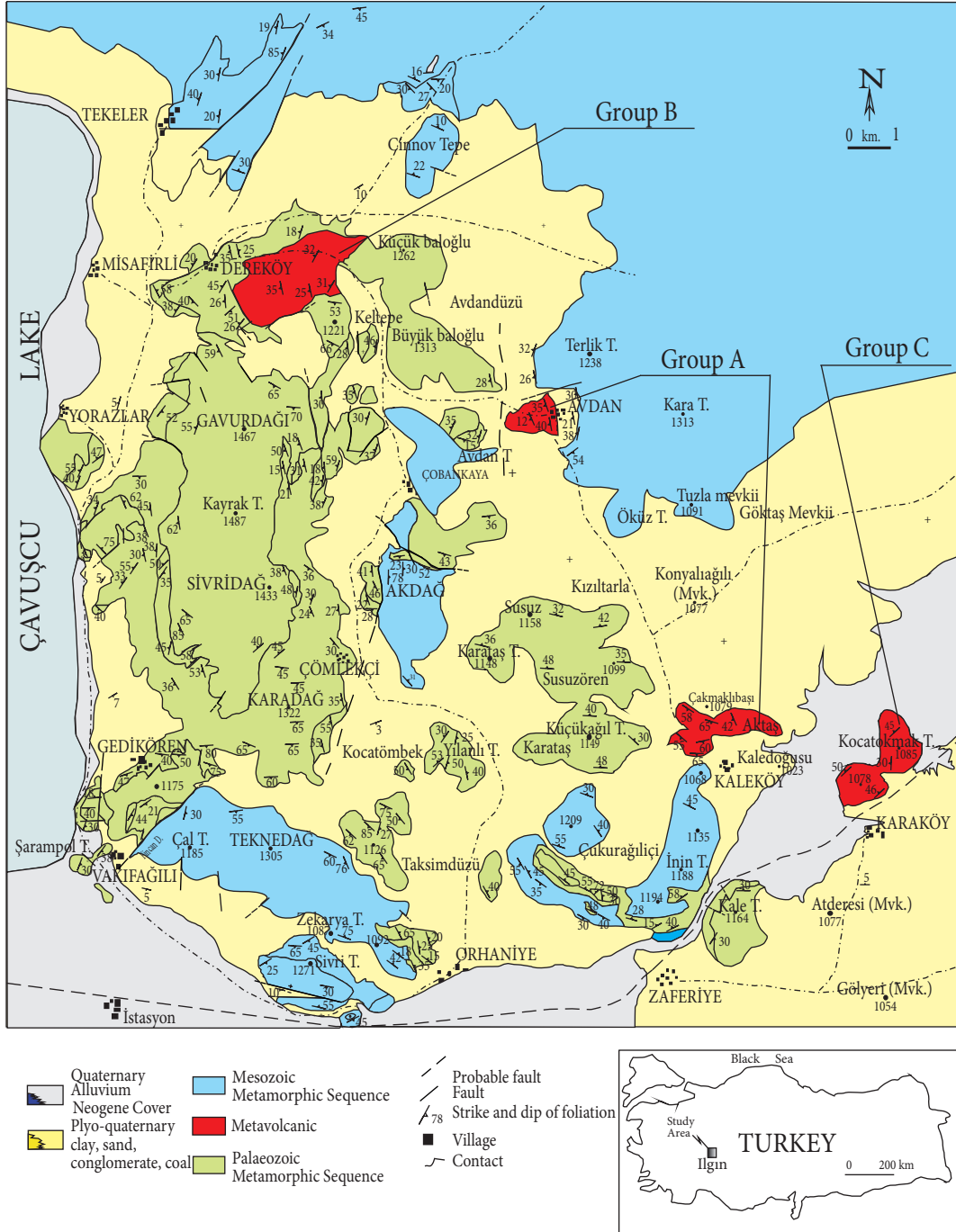


Figure 2. Location and geology map of the study area (from Özdamar et al. 2012).

$$\gamma_d = (1 - n)\gamma_s \quad (1)$$

or

$$\gamma_d = \frac{\gamma_s}{1 + e} \quad (2)$$

where  $n$  is porosity and  $e$  is void ratio.

Specific gravity is calculated by inserting measured

values of dry unit weight, porosity, and void ratio. The analysed specific gravity of rock mass ( $\gamma_{s(a)}$ ) is measured with a pycnometer. The estimated specific gravity ( $\gamma_{s(c)}$ ) is calculated by using Eqs. (1) or (2). The  $RCV_p$  is generated from the following conditions:

1. If the analysed specific gravity of rock mass ( $\gamma_{s(a)}$ ) is greater than the calculated specific gravity of rock mass ( $\gamma_{s(c)}$ ), then:

**Table 1.** Geochemical analyses of the metarhyolites (data from Özdamar *et al.* 2012).

Sample no.	SiO <sub>2</sub> (%)	Al <sub>2</sub> O <sub>3</sub> (%)	Fe <sub>2</sub> O <sub>3</sub> (%)	CaO (%)	MgO (%)	Na <sub>2</sub> O (%)	K <sub>2</sub> O (%)	TiO <sub>2</sub> (%)	MnO (%)	P <sub>2</sub> O <sub>5</sub> (%)	Ba (ppm)	Zr (ppm)	LOI (%)
A3	70	17.69	1.55	0.1	1.11	0.08	6.64	0.05	0.01	0.05	100	61	2.7
A1	74.87	13.13	1.45	0.45	0.13	1.77	5.8	0.08	0.01	0.06	61	87	1.1
A2	74.58	12.85	0.93	0.94	0.17	0.51	8.09	0.08	0.01	0.04	106	92	1.8
A4	72.24	16.29	1.72	0.06	0.15	0.21	7.82	0.009	0.01	0.03	74	35	1.5
B1	66.91	17.05	1.8	0.9	0.42	0.22	9	0.37	0.02	0.07	134	53	2.8
B2	76.92	12.91	1.2	0.11	0.46	0.07	6.21	0.08	0.009	0.009	79	205	1.5
B3	67.39	17.16	1.66	0.14	0.37	0.14	10.66	0.37	0.01	0.06	299	518	1.8
C1	74.83	13.52	1.05	0.19	0.08	0.76	8.08	0.02	0.009	0.03	177	69	1.4
C2	72.20	14.28	1.46	0.08	0.15	0.26	9.42	0.03	0.02	0.02	139	30.5	1.4

$$RCV_p = \frac{\gamma_{s(c)}}{\gamma_{s(a)}} \quad (3)$$

2. If the analysed specific gravity of rock mass ( $\gamma_{s(a)}$ ) is less than the calculated specific gravity of rock mass ( $\gamma_{s(c)}$ ), then:

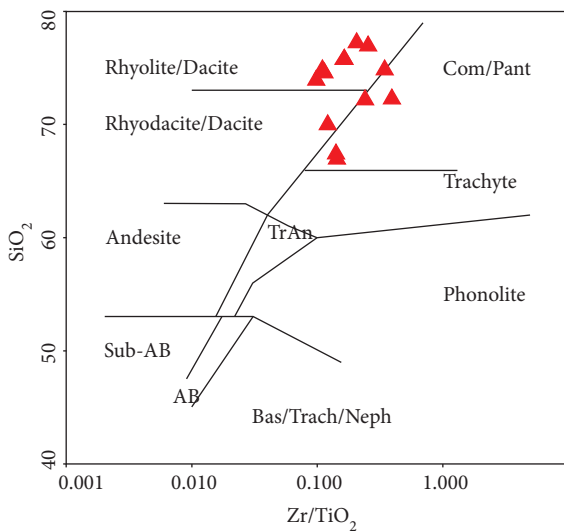
$$RCV_p = \left| 1 + \left[ 1 - \frac{\gamma_{s(c)}}{\gamma_{s(a)}} \right] \right| \quad (4)$$

Eqs. (3) and (4) introduce a new method for quantifying the degree of weathering and alteration in rock material based on specific gravity. Eq. (4) is derived from Eq. (3) empirically. The resultant  $RCV_p$  value ranges from 1 to 0, where the  $RCV_p$  values of fresh rocks are equal to or very close to 1. In intensively weathered and altered

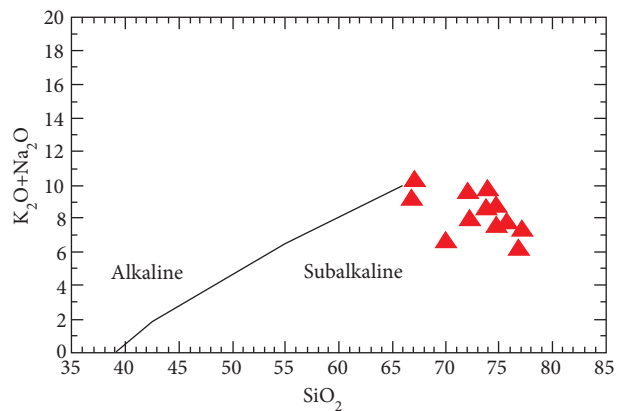
rock processes,  $RCV_p$  values approach 0. The  $RCR_p$  is also defined with the changes in the entire volume of rocks, as shown in Eq. 5.

$$RCR_p = \left[ \frac{1 - RCV_p}{1 - (1 - RCV_p)} \right] \times 100 \quad (5)$$

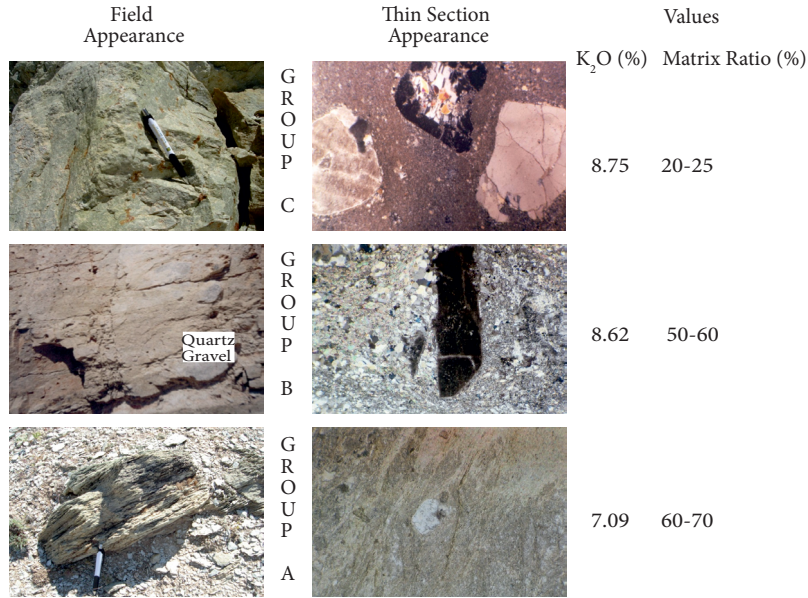
In  $RCR_p$ , the fresh rock volume is 1, and according to the phase diagram of the soil or rock, the  $RCV_p$  shows the present status of rock in the field. The numerator in Eq. (5) represents the changed portion of the rock volume (eroded material) and the denominator is the residual part of the rock volume. As with  $RCV_p$ ,  $RCR_p$  shows the degree of weathering or chemical alteration.  $RCR_p$  values near 100% indicate strong weathering and chemical alteration whereas values near 0% indicate minimal weathering and alteration effects. For values of  $RCV_p$  equal to or less than 0.5, Eq. (5) provides values exceeding 100%, a condition impossible in nature.



**Figure 3.** Classification of the metarhyolites of the Iğın area using the  $Zr/TiO_2$ -Nb/Y diagram of Winchester & Floyd (1977) (from Özdamar *et al.* 2012).



**Figure 4.** Compositions of the Iğın metavolcanics using the  $SiO_2$ - $Na_2O+K_2O$  diagram of Irvine & Baragar (1971) (from Özdamar 2011).



**Figure 5.** Field and photomicrographs of metarhyolites.

The metarhyolites are classified in 3 groups, namely A, B, and C, which represent their weathering degree from weathered to fresh rocks in terms of their  $RCV_p$  values. The appearance of the metarhyolites and their representative thin section images are given in Figure 5, and the  $RCV_p$  and  $RCR_p$  values of the 3 main groups and their subgroups in the study area are given in Table 2.

In Table 2, the  $RCV_p$  and  $RCR_p$  values are taken from 5 samples from each subgroup and the 3 main groups'  $RCV_p$  and  $RCR_p$ . Values are the average values of the subgroups for each main group, and these values are very approximate in Figure 5. The  $RCV_p$  and  $RCR_p$  values are given in order from fresh to changed (modified) rock in terms of the average  $RCV_p$  and  $RCR_p$  values of the 3 main groups in Table 1.

The  $RCV_p$  values show that the rocks in group C are fresh and the following 2 main groups may be classified as

group B, which is moderately altered, and group A, which is altered. The relationship between  $RCV_p$  and  $RCR_p$  can be described by the following formula, and the correlation coefficient of 0.9999 is given in Figure 6.

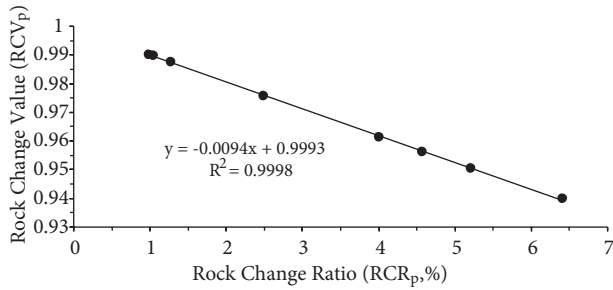
$$RCV_p = -(0.0094 \times RCR_p) + 0.9993 \quad (6)$$

#### 4. Relationship between $RCV_p$ - $RCR_p$ and physico-mechanical properties

Physical properties not only describe the present conditions of the rocks but are also used as a useful tool for describing engineering properties. The physical properties of the samples studied were analysed based on Turkish Standard TS-699 (Turkish Standards Institution 1982). The average values of physical properties from the study area are given in Table 3.

**Table 2.** The  $RCV_p$  and  $RCR_p$  values of the metarhyolites.

Group name		$RCV_p$	$RCR_p$	Average $RCV_p$	Average $RCR_p$
C	C1	0.990203	0.989379	0.990203	0.989379
	B1	0.989678	1.042992		
B	B2	0.987389	1.277255	0.984221	1.492455
	B3	0.975653	2.495427		
A	A4	0.961439	4.010737	0.951973	4.972632
	A3	0.956346	4.56472		
	A2	0.950466	5.2116		
	A1	0.939777	6.408197		



**Figure 6.** The relationship between RCV<sub>p</sub> and RCR<sub>p</sub> in the study area.

Density and porosity are 2 fundamental properties of rocks. Density is influenced primarily by both mineral composition and void space, where increasing void space increases porosity and decreases density. The IAEG Commission (1979) grouped the dry density and porosity of rocks into 5 classes, as shown in Table 4.

The rocks studied in the Ilgın area have moderate to high dry density with medium to very low porosity. This description shows that the rock weathering has been affected by atmospheric conditions and that clay products fill the void spaces and cracks in the rocks. The relationships between RCR<sub>p</sub> values and physical properties are polynomial (Figure 7, Table 5).

Eqs. (7) through (11), given in Table 5, show that rocks in the study area have a specific gravity ( $\gamma_s$ ) of 3244.50 kg/m<sup>3</sup> and a dry unit weight ( $\gamma_d$ ) of 3222.30 kg/m<sup>3</sup> when they start to interact with atmospheric conditions (fresh rock) by uplift after metamorphism, in which the RCR<sub>p</sub> value is equal to 0% and RCV<sub>p</sub> is equal to 1. These equations also explain that rocks in the study area were in the monolith phase and subsequently cracks developed inside rocks, producing porosity (0.0485%) with a 0.75% RCR<sub>p</sub> and void ratio (0.00075) with a 0.77% RCR<sub>p</sub>, and atmospheric water was held (0.0206%) at a RCR<sub>p</sub> value of 0.79%.

Moreover, after the development of porosity with a RCR<sub>p</sub> value of 0.75% and a RCV<sub>p</sub> value of 0.99225 in the

area, the specific gravity ( $\gamma_s$ ) value of 2766.28 kg/m<sup>3</sup> and dry unit weight ( $\gamma_d$ ) value of 2731.42 kg/m<sup>3</sup> imply that the decrease of these values from the monolith phase is 14.74% for specific gravity and 15.23% for dry unit weight (the RCV<sub>p</sub> changed by only 0.00775). The change of physical conditions shows that the weathering rate was fast with very small RCV<sub>p</sub> values. These equations (Table 5) explain that the rocks in the study area will change completely at a RCR<sub>p</sub> value of 9.01%, at which the RCV<sub>p</sub>, specific gravity ( $\gamma_s$ ), dry unit weight ( $\gamma_d$ ), water content ( $w$ ), porosity ( $n$ ), and void ratio ( $e$ ) values would be equal to 0.914606, 278.33 kg/m<sup>3</sup>, 142.76 kg/m<sup>3</sup>, 13.25%, 29.75%, and 0.35, respectively. In these conditions, the rocks remain as metarhyolite chemically, but their physico-mechanical properties and mineralogical compositions will continue to change and the metarhyolites will become soil.

Mechanical properties of the rocks in the study area were investigated by point load tests based on those of the International Society for Rock Mechanics (ISRM 1985). The average values of the point load strength index of each sub- and main sample group in the study area are given in Table 6.

One of the useful and commonly used strength classifications for rocks is the point load strength index, shown in Table 7, devised by Franklin & Broch (1972).

The rocks in the study area vary from medium to extremely high strength according to the observed degree of weathering, which increases from group C to group A. However, the average values are in a very high to extremely high strength class, in which the group C rocks represent fresh samples with a 0.990 RCV<sub>p</sub> and groups B and A represent the moderate weathering development samples with 0.984 and 0.952 RCV<sub>p</sub> values, respectively. The relationship between RCR<sub>p</sub> values versus point load strength index values also exhibits a polynomial relationship, with a correlation coefficient of 0.5667 defined in Eq. (12) and shown in Figure 8.

According to Eq. (12), rock strength was 27.08 MPa when the rocks were in the monolith phase, and the

$$I_{s(50)} = -0.3109x(RCR_p)^3 + 4.4355(RCR_p)^2 - 17.726(RCR_p) + 27.08 \quad (12)$$

**Table 3.** Physical properties of the metarhyolites.

Group name	Specific gravity ( $\gamma_s$ ), kg/m <sup>3</sup>	Dry unit weight ( $\gamma_d$ ), kg/m <sup>3</sup>	Saturated unit weight ( $\gamma_{sat}$ ), kg/m <sup>3</sup>	Water content ( $w$ ), %	Porosity ( $n$ ), %	Void ratio ( $e$ )	
A	A1	2777.78	2570.34	2586.51	0.63	1.61	0.02
	A2	2777.78	2624.49	2631.23	0.25	0.67	0.01
	A3	2922.51	2721.47	2746.96	0.92	2.51	0.03
	A4	2764.50	2630.70	2640.74	0.37	0.98	0.01
B	B1	2591.70	2549.55	2575.19	1.00	2.55	0.03
	B2	2672.80	2646.22	2668.87	0.85	2.25	0.02
	B3	2416.99	2323.03	2385.97	2.45	5.70	0.06
C	2665.76	2608.78	2620.55	0.45	1.17	0.01	

**Table 4.** Dry density and porosity classification of rocks (IAEG Commission 1979).

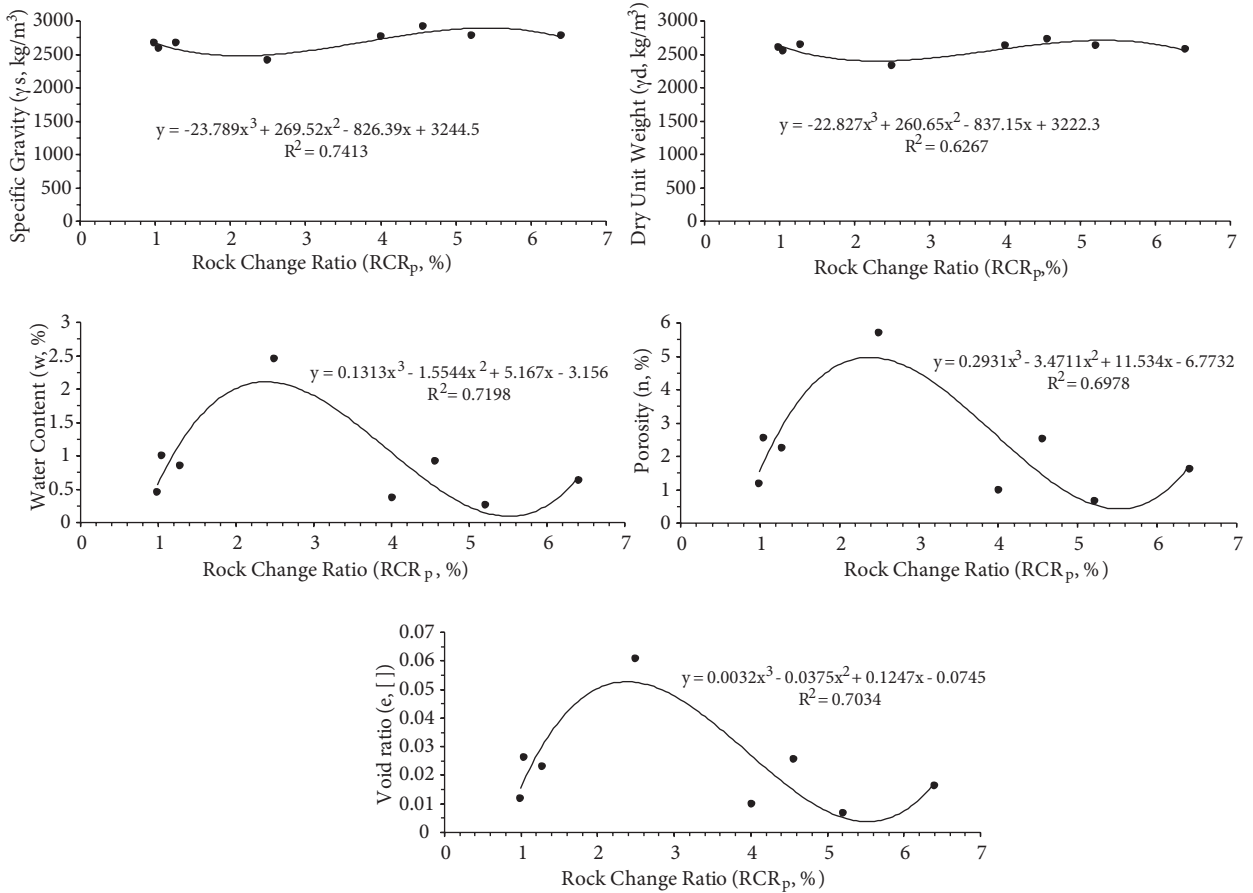
Class	Dry density (Mg/m <sup>3</sup> )	Description	Porosity (%)	Description
1	Less than 1.8	Very low	Over 30	Very high
2	1.8-2.2	Low	15-30	High
3	2.2-2.55	Moderate	5-15	Medium
4	2.55-2.75	High	1-5	Low
5	Over 2.75	Very high	Less than 1	Very low

changed rock strengths were 16.15 MPa for 0.75% RCR<sub>p</sub> (RCV<sub>p</sub> = 0.99225), 15.92 MPa for 0.77% RCR<sub>p</sub> (RCV<sub>p</sub> = 0.992062), 15.69 MPa for 0.79% RCR<sub>p</sub> (RCV<sub>p</sub> = 0.991874), and 0.041 MPa for 9.01% RCR<sub>p</sub> (RCV<sub>p</sub> = 0.914606). These changes in the rock strength values also indicate high weathering rates.

The K-Ar ages of the metarhyolites exhibit a meaningful relation between RCV<sub>p</sub> and RCR<sub>p</sub> values. The K-Ar age is

60.4 ± 0.9 Ma for group C, 62.6 ± 0.9 Ma for group B, and 64.1 ± 2.00 Ma for group A. The relationships between K-Ar ages and RCV<sub>p</sub> and RCR<sub>p</sub> values are given in Figure 9 and Table 8.

Eqs. (13) through (18), given in Table 8, make a good approach to full-rock changed age. The calculated results for the estimation of the full lifetime of rock change are given in Table 9. The remaining rock lifetime can



**Figure 7.** Relationships between RCR<sub>p</sub> and physical properties of the metarhyolites.

**Table 5.** Relationships between  $RCR_p$  and physical properties of the metarhyolites.

Relationships between $RCR_p$ and physical properties		Correlation coefficient (r)
$\gamma_s = -23.789 \times (RCR_p)^3 + 269.52 \times (RCR_p)^2 - 826.39 \times (RCR_p) + 3244.5$	(7)	0.861
$w = 0.1313 \times (RCR_p)^3 - 1.5544 \times (RCR_p)^2 + 5.167 \times (RCR_p) - 3.156$	(8)	0.848
$e = 0.0032 \times (RCR_p)^3 - 0.0375 \times (RCR_p)^2 + 0.1247 \times (RCR_p) - 0.0745$	(9)	0.839
$n = 0.2931 \times (RCR_p)^3 - 3.4711 \times (RCR_p)^2 + 11.534 \times (RCR_p) - 6.7732$	(10)	0.835
$\gamma_d = -22.827 \times (RCR_p)^3 + 260.65 \times (RCR_p)^2 - 837.15 \times (RCR_p) + 3222.3$	(11)	0.792

**Table 6.** Point load strength index values of the metarhyolites.

Group name	Point load strength index values, $I_{s(50)}$ (MPa)	Point load strength index average values, $I_{s(50)}$ (MPa)
A	A1	13.77
	A2	13.93
	A3	0.99
	A4	13.23
B	B1	11.31
	B2	8.04
	B3	5.76
C	17.55	17.55

**Table 7.** Point load strength classification (Franklin & Broch 1972).

Description	Point load strength index (MPa)	Equivalent uniaxial compressive strength (MPa)
Extremely high strength	Over 10	Over 160
Very high strength	3-10	50-160
High strength	1-3	15-60
Medium strength	0.3-1	5-16
Low strength	0.1-0.3	1.6-5
Very low strength	0.03-0.1	0.5-1.6
Extremely low strength	Less than 0.03	Less than 0.5

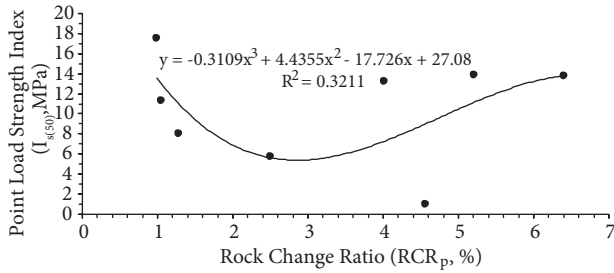
be calculated simply by making a subtraction between current and calculated ages of rocks. These results are given in Table 10.

The highest correlation coefficient values were gathered for the maximum K-Ar age values versus  $RCV_p$  and  $RCR_p$  values ( $r = 0.94$ ). The evaluations show that the whole-rock change time by weathering, disintegration, and maybe alteration will be complete after 4.58 Ma for highly weathered rocks (group A), 7.18 Ma for moderately

weathered rocks (group B), and 9.38 Ma for fresh or weakly weathered rocks (group C), respectively. These values are geometric means of the calculated minimum, average, and maximum age values of each group.

The group A rocks were formed  $230 \pm 2$  Ma ago and group B rocks  $229 \pm 2$  Ma ago (Özdamar 2011). Metamorphism occurred at  $63.73 \pm 0.06$  Ma in group A rocks and  $62.64 \pm 0.12$  Ma in group C rocks (Özdamar 2011). The relationships between formation ages (FAs) of





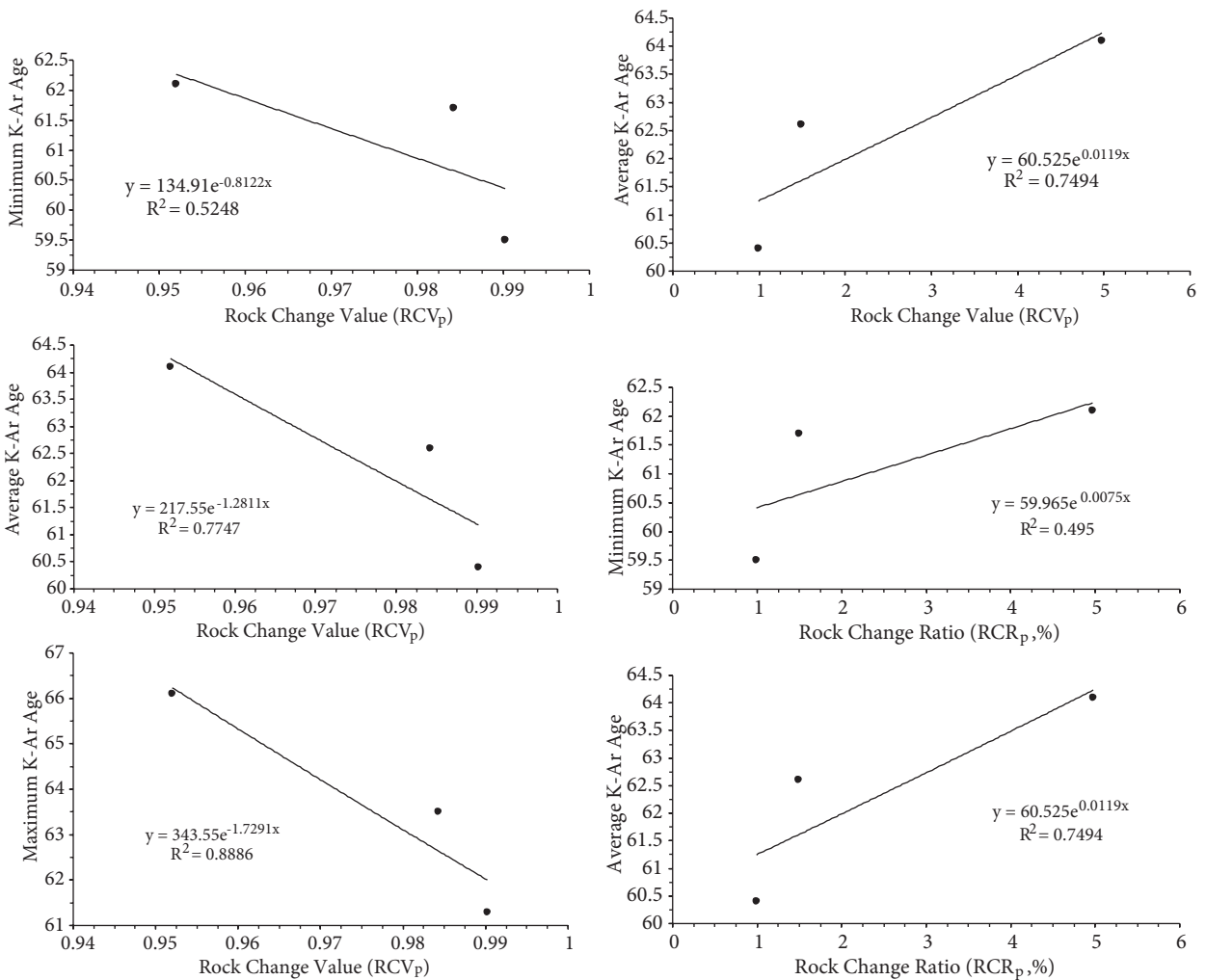
**Figure 8.** Relationship between  $RCR_p$  and point load strength index of the rocks in the study area.

minimum, average, and maximum values and  $RCV_p$  and  $RCR_p$  values are given in Figure 10 and Table 11. Eqs. (19) through (24), shown in Table 11, give the zone C rocks a formation age of  $228 \pm 2$  Ma. The results also explain why the rocks of this zone remain fresh.

The relationships between metamorphism ages (MtAs) of minimum, average, and maximum values and  $RCV_p$  and  $RCR_p$  values are given in Figure 11 and Table 12. Eqs. (25) through (30) give the metamorphism age of  $62.68 \pm 0.11$  Ma for group B rocks. This age is close to the metamorphism age of group C and explains the moderate rock-change conditions in group B rocks shown in Figure 5.

**5. Results and conclusion**

The metamorphic sequences in the Ilgin area have rocks of both Palaeozoic and Mesozoic ages. Metavolcanics are subalkaline and range from rhyodacite to rhyolite. These rocks have moderate to high dry density versus medium to very low porosity values, and their strengths change from medium to extremely high values. With these characteristics, the rock weathering in the field can be determined from fresh to moderately changed conditions.



**Figure 9.** Relationships between K-Ar ages and  $RCV_p$  -  $RCR_p$  values of the metarhyolites.

**Table 8.** Relationships between K-Ar ages and  $RCV_p$  and  $RCR_p$  values of the metarhyolites.

K-Ar ages	Relationships between K-Ar age and $RCV_p$	Correlation coefficient (r)
Minimum	$(K-Ar)_{min} = 134.91 \times e^{-0.8122 \times (RCV_p)}$ (13)	0.724
Average	$(K-Ar)_{ave} = 217.55 \times e^{-1.2811 \times (RCV_p)}$ (14)	0.880
Maximum	$(K-Ar)_{max} = 343.55 \times e^{-1.7291 \times (RCV_p)}$ (15)	0.943
K-Ar ages	Relationships between K-Ar age and $RCR_p$	Correlation coefficient (r)
Minimum	$(K-Ar)_{min} = 59.965 \times e^{0.0075 \times (RCR_p)}$ (16)	0.704
Average	$(K-Ar)_{ave} = 60.525 \times e^{0.0119 \times (RCR_p)}$ (17)	0.866
Maximum	$(K-Ar)_{max} = 61.092 \times e^{0.0162 \times (RCR_p)}$ (18)	0.932

**Table 9.** Lifetime of complete rock changes: calculations between K-Ar ages and  $RCV_p$  and  $RCR_p$  values.

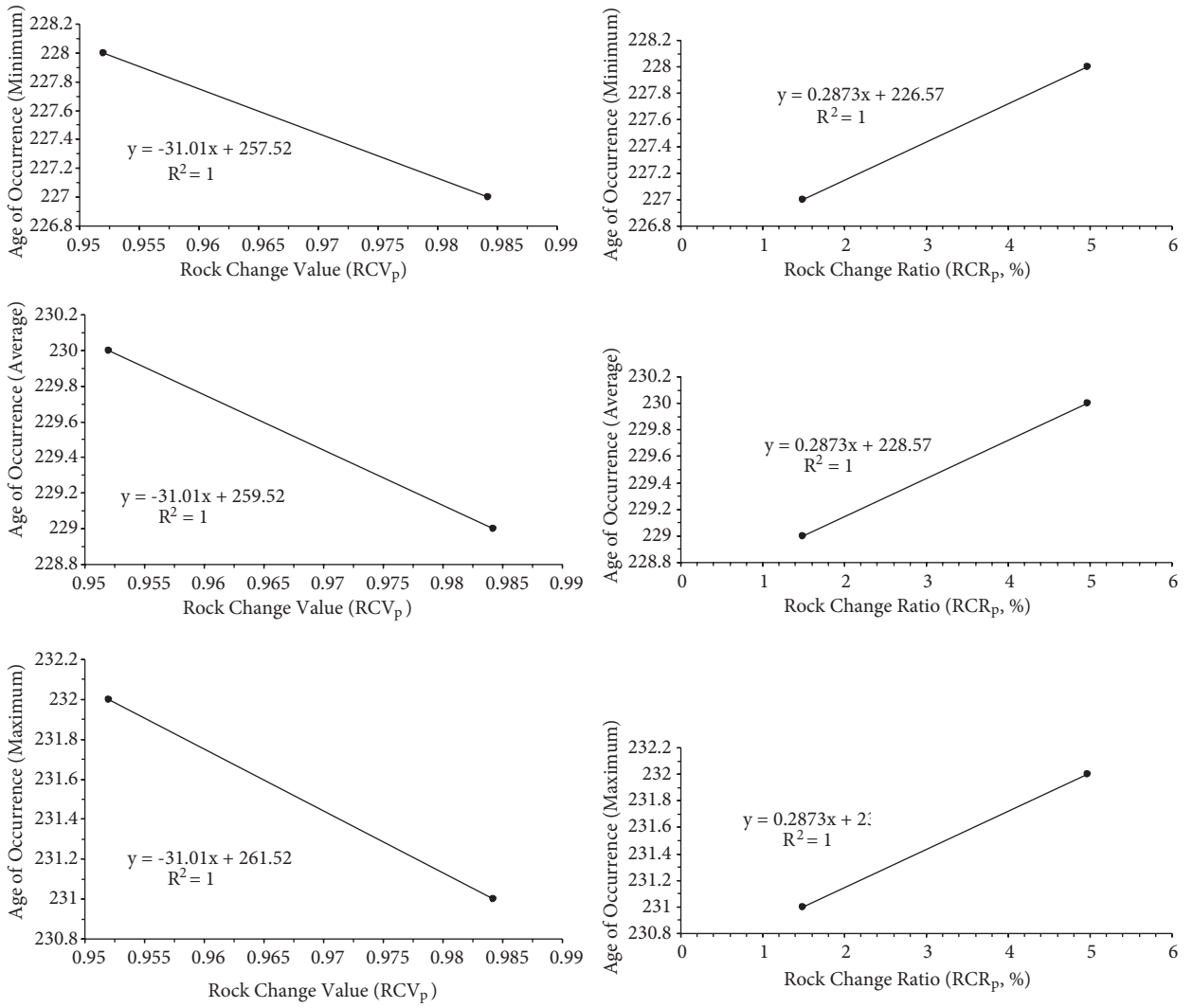
K-Ar ages according to $RCV_p$ (Ma)			K-Ar ages according to $RCR_p$ (Ma)		
Minimum (Ma)	Average (Ma)	Maximum (Ma)	Minimum (Ma)	Average (Ma)	Maximum (Ma)
64.184	67.405	70.660	64.157	67.375	70.693

**Table 10.** K-Ar age calculations according to  $RCV_p$  and  $RCR_p$  values of the metarhyolites.

Group name	Remaining time according to the relationship between K-Ar ages and $RCV_p$ values for fully changed rock			Remaining time according to the relationship between K-Ar ages and $RCR_p$ values for fully changed rock		
	Minimum (Ma)	Average (Ma)	Maximum (Ma)	Minimum (Ma)	Average (Ma)	Maximum (Ma)
C	4.684	7.005	9.360	4.657	6.975	9.393
B	2.484	4.805	7.160	2.457	4.775	7.193
A	2.084	3.305	4.560	2.057	3.275	4.593

In fact, the change of physical conditions of the rocks helps us to understand the high speed of the weathering with very small  $RCV_p$  values. Additionally, these rocks were divided into 3 groups defined by their weathering conditions with  $RCV_p$  and  $RCR_p$  values in order to define the primary physical conditions of the rocks using the relationships between  $RCR_p$  and physical properties (Figure 7, Table 5). These equations show that the strength

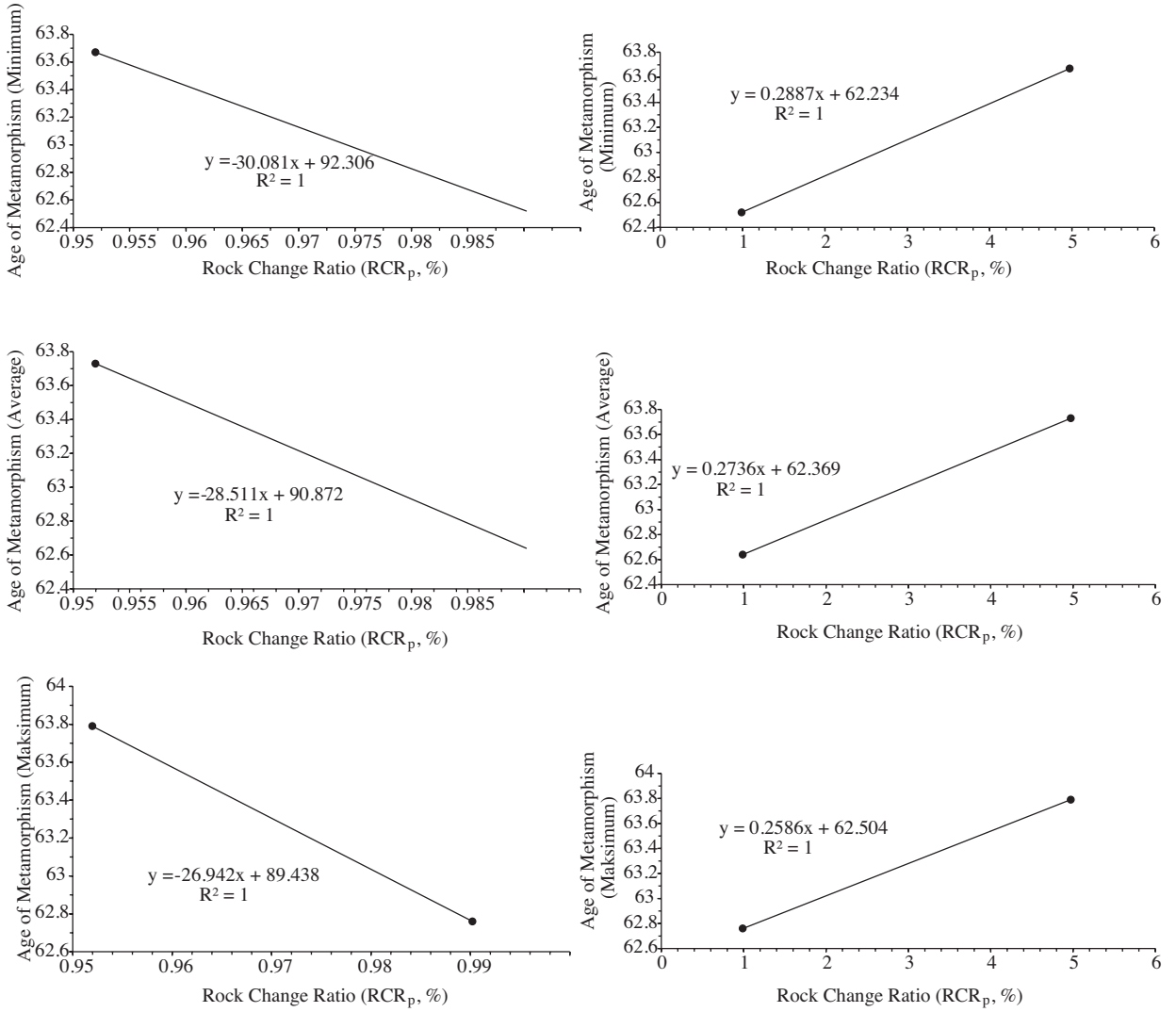
in the monolith-phase rocks was 27.08 MPa, specific gravity ( $\gamma_s$ ) was 3244.50 kg/m<sup>3</sup>, and dry unit weight ( $\gamma_d$ ) was 3222.30 kg/m<sup>3</sup>, where the  $RCR_p$  value is equal to 0% and  $RCV_p$  is equal to 1. After the development of fractures and cracking, the primary porosity,  $n$ , became 0.0485%, with a  $RCR_p$  value of 0.75%. The voids in bulk composition developed with a 0.77%  $RCR_p$  value where the void ratio is  $e = 0.00075$ , and atmospheric water would have been held



**Figure 10.** Relationships between formation ages (FAs) and RCV<sub>p</sub> - RCR<sub>p</sub> values of the rocks in the study area.

**Table 11.** Relationships between formation ages (FAs) and RCV<sub>p</sub> and RCR<sub>p</sub> values of the rocks in the study area.

FAs	Relationships between FA and RCV <sub>p</sub>	Correlation coefficient (r)
Minimum	$(FA)_{\min} = -31.01 \times RCV_p + 257.52$ (19)	1
Average	$(FA)_{\text{ave}} = -31.01 \times RCV_p + 259.52$ (20)	1
Maximum	$(FA)_{\max} = -31.01 \times RCV_p + 261.52$ (21)	1
FAs	Relationships between FA and RCR <sub>p</sub>	Correlation coefficient (r)
Minimum	$(FA)_{\min} = 0.2873 \times RCR_p + 226.57$ (22)	1
Average	$(FA)_{\text{ave}} = 0.2873 \times RCR_p + 228.57$ (23)	1
Maximum	$(FA)_{\max} = 0.2873 \times RCR_p + 230.57$ (24)	1



**Figure 11.** Relationships between metamorphism ages (MA) and  $RCV_p$  -  $RCR_p$  values of the rocks in the study area.

**Table 12.** Relationships between timing of metamorphism (MA) and  $RCV_p$  and  $RCR_p$  values of the rocks in the study area.

MA ages	Relationships between MtA age and $RCV_p$	Correlation coefficient (r)
Minimum	$(MtA)_{min} = -30.081 \times RCV_p + 92.306$ (25)	1
Average	$(MtA)_{ave} = -28.511 \times RCV_p + 90.872$ (26)	1
Maximum	$(MtA)_{max} = -26.942 \times RCV_p + 89.438$ (27)	1
MA ages	Relationships between MtA age and $RCR_p$	Correlation coefficient (r)
Minimum	$(MtA)_{min} = 0.2887 \times RCR_p + 62.234$ (28)	1
Average	$(MtA)_{ave} = 0.2736 \times RCR_p + 62.369$ (29)	1
Maximum	$(MtA)_{max} = 0.2586 \times RCR_p + 62.504$ (30)	1

( $w = 0.0206\%$ ), for the  $R_{CR_p}$  is equal to 0.79%. The  $R_{CV_p}$  and  $R_{CR_p}$  values also show that rocks in the study area change completely with the 9.01%  $R_{CR_p}$  value. Dry unit weight will also be smaller than  $1.4 \text{ Mg/m}^3$  when the  $R_{CR_p}$  is less than 8.11%. In this case, the Ilgın area rocks will also remain as rhyolite whereas the physical, mechanical, and mineralogical properties of the rocks will change for the 9.01%  $R_{CR_p}$  value and the soil structure will have formed. The  $R_{CV_p}$  value, strength, specific gravity ( $\gamma_s$ ), dry unit weight ( $\gamma_d$ ), water content ( $w$ ), porosity ( $n$ ), and void ratio ( $e$ ) values will be equal to 0.914606, 0.041 MPa, 278.33  $\text{kg/m}^3$ , 142.76  $\text{kg/m}^3$ , 13.25%, 29.75%, and 0.35, respectively, for the 9.01%  $R_{CR_p}$ .

The rock change was compared with the K-Ar age values rather than the formation ages due to the new

formation of the rocks after metamorphism, where the weathering from fresh to weathered rock starts again. The full-rock change is related to the alkali ratio of the rocks. In the Ilgın area the alkali ratio of the rocks changes between 8.12% and 9.40% (the average is 8.89%), where the  $R_{CR_p}$  is 9.01% when the total alteration of all rocks will occur. The correspondence between these results explains that the  $R_{CV_p}$  and  $R_{CR_p}$  values are also very useful index values for determining both the initial and current situation of the rock easily. Furthermore, they also give an opportunity to predict the future condition of the rocks. In this study, the rock conditions in the Ilgın area were forecast with the aim of comparing the K-Ar age values versus  $R_{CV_p}$  and  $R_{CR_p}$  values. The interpretations show that the timing of full rock alteration in the Ilgın area is between 4.6 and 9.4 Ma.

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